Decline in groundwater levels in Thang Long Industrial Park within the area of Northwest Hanoi, Vietnam

Noboru HIDA*1 and Nguyen Van GIANG*2

Abstract

The purpose of this paper is to urgently report the decline in groundwater levels at Thang Long Industrial Park (TLIP) in the area of Northwest Hanoi, Vietnam. The total pumping capacity in TLIP is 56,970 m³/day. According to the records of the observation well close to TLIP, groundwater levels have declined since 2008. The amount of decline was approximately 5.67m between 2008 and 2011. Lower levels appear in the dry season, particularly in March-April. In order to maintain sustainable use of groundwater in TLIP, one of the issues is to carry out an artificial recharge of groundwater using Kim Hoa Canal, of which water is supplied by the Ca Lo River.

Key words: groundwater levels, artificial recharge of groundwater, Thang Long Industrial Park, Hanoi

I. Introduction

Hanoi: 21°02’00"N 105°51’00"E, capital of Vietnam, has an area of 3,344 square kilometers and a population of 6,500,000 in 2010. Downtown of Hanoi is developed on the right bank of the Red River (Song Hong). In and around the downtown area, domestic and industrial water supplies have been relying on groundwater. A public water supply plant using groundwater was constructed first in 1909 with a total pumping capacity of 15,000 m³/day. At present, the total yield is approximately 700,000 m³/day, made up as follows: 10 public water supply plants 420,000 m³/day, relatively big water users such as schools and hospitals 150,000 m³/day and private household 150,000 m³/day, respectively (Dan and Dzung, 2002; Khanh et al., 2011, p.5). In the process of the increase, groundwater problems have occurred such as decline of the water table (Fischer et al., 2011), land subsidence (Nguyen and Helm, 1995; Thu and Fredlund, 2000) and groundwater pollution (Berg et al., 2001; Dan and Dzung, 2002).

On the other hand, modern industrial zones have been rapidly established in recent years on the left bank of the Red River, which was once rice paddy fields. In these zones, one of the most significant developments is Thang Long Industrial Park (TLIP), 2.75 km², located in the northwest of the downtown area. Some factories have been in operation since 2007.

In TLIP, water for industry and drinking is drawn from underground. However, groundwater levels have already declined for the last several years. The purpose of this paper is to urgently report the decline based on observed records. The intention is to explore the way of sustainable use of groundwater in the future, avoiding some similar groundwater problems caused in the right bank area of the Red River. Some results of hydrological and hydrogeological conditions including TLIP have already been published (Giang and Hida 2009).

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and 2010; Hida et al., 2010).

II. Study area

II -1 Site

The site of the study area, within Hanoi City, between downtown and Noi Bai International Airport, is shown in Fig. 1. In this area, the Red River (on the south side) and Ca Lo River (on the north side) flow towards the east.

Fig. 1 Study area in Northwest Hanoi.
Legend: TLIP: Thang Long Industrial Park, OW: Observation well 1,2,3, PWG: Pumping well gallery (west and east), WT: Water treatment plant, QMIZ: Quang Minh Industry Zone (Base map is by the Topographic Map HANOI and SON TAY (1:50,000), 2001, Hanoi Vietnam.)

II -2 Thang Long Industrial Park (TLIP)

TLIP bordering on the Red River is an important industrial zone in Hanoi, which was completely reclaimed in 2007. Industrial water sources are limited to groundwater. In addition, Quang Minh Industrial Zone (QMIZ) exists in the north, where groundwater yield is small.

II -3 Groundwater yield in TLIP

Pumping well gallery (PWG)-west and PWG-east are located on the west side of TLIP (Fig. 1). Eight pumping wells each have been constructed in the PWG-west (inside the river bank) and in the PWG-east (outside the river bank). Wells in the PWG-west are 57 - 66m deep, each with pumping capacity of 4,320 m³/day. Total pumping capacity of the eight wells is 34,560 m³/day. Wells in PWG-east are 62 - 67m deep. The pumping capacity of each is 2,800 m³/day, and the total for the eight wells is 22,400 m³/day. The total pumping capacity of all 16 wells is 56,970 m³/day. All of the groundwater is withdrawn from the Pleistocene aquifer mentioned later.

II -4 Physical environment

Climate: Hanoi belongs to the monsoon region. A hydrological year consists of a rainy season from May to October, and a dry season from November to April.

Topography: The study area is a floodplain formed by the Red River and Ca Lo River (DGMV, 2005). The ground surface is almost 10m above sea level and inclined from the side of the Ca Lo River towards the side of the Red River approximately in $1.0 \times 10^2$ (HXBBD, 2001).

Hydrogeology: Most of the groundwater in the Red River Delta including the study area is contained in Quaternary sediments (Lan et al., 2007). The aquifer is composed of two layers, the upper Holocene aquifer (Qh) and the lower Pleistocene aquifer (Qp). The thickness of Qh is estimated to be 10-15m and that of Qp is about 20-40m (Dan and Dzung, 2002; DGMV, 2005). Characteristic properties of these two in the Red River Delta have been discussed in detail (Bui et al., 2012; Giang et al., 2012; Jusseret et al., 2010; Hori et al., 2004).

III. Observation wells

Three observation wells: OW1, OW2 and OW3 were established in the study area. OW1 is substantially situated in the middle of PWG-west and PWG-east of TLIP (Fig. 1). OW2 and OW3 were placed to the north of OW1. Specifications and geologic
Table 1 Specifications of the OW1, OW2 and OW3.

<table>
<thead>
<tr>
<th>Items</th>
<th>OW1</th>
<th>OW2</th>
<th>OW3</th>
</tr>
</thead>
<tbody>
<tr>
<td>Site (See Fig. 1)</td>
<td>105°54’20&quot; E</td>
<td>105°45’00&quot; E</td>
<td>105°44’05&quot; E</td>
</tr>
<tr>
<td></td>
<td>21°06’57&quot; N</td>
<td>21°12’20&quot; N</td>
<td>21°10’24&quot; N</td>
</tr>
<tr>
<td>Ground level (Elv. m)</td>
<td>10.0</td>
<td>10.0</td>
<td>8.0</td>
</tr>
<tr>
<td>Well head (Elv. m)</td>
<td>10.0</td>
<td>11.2</td>
<td>8.0</td>
</tr>
<tr>
<td>Depth (m)</td>
<td>60.0</td>
<td>48.0</td>
<td>40.0</td>
</tr>
<tr>
<td>Screen length (m)</td>
<td>8.0</td>
<td>4.0</td>
<td>4.0</td>
</tr>
<tr>
<td>Diameter, (mm)</td>
<td>90</td>
<td>90</td>
<td>90</td>
</tr>
<tr>
<td>Logger Recorded by logger</td>
<td>Water level and Water temperature from July 2008</td>
<td>See Fig. 2</td>
<td></td>
</tr>
<tr>
<td>Construction</td>
<td>March 2008</td>
<td>June 2008</td>
<td>July 2008</td>
</tr>
<tr>
<td>GCS(3)</td>
<td></td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

Notes: (1) Drilling depth from ground level (m) (2) Installed at the bottom of the well, See Fig. 2
(3) Geologic columnar section

columnar section of the three wells are shown in Table 1 and Fig. 2. A screen installed in OW1 is in the layer of the lower Pleistocene aquifer (Qp).

A logger to record groundwater level and temperature was inserted to the depth of 15m from the surface in each well. The recording interval of each was one hour.

IV. Results: decline in groundwater levels at OW1

Annual change in groundwater levels (Elevation m: Elv. m) at OW1 in Thang Long Industrial Park is shown in Fig. 3 which is based on the observation records of daily mean from 2008 to 2011 hydrological year (hy). Table 2 shows average groundwater levels (Elv. m) at OW1 by hydrological year, by rainy and dry seasons, and by month in the dry season. Fig. 3 and Table 2 do not include the records for April 2011 hy due to the lack of records. The findings based on Fig. 3 and Table 2 are as follows.

Firstly, groundwater level has been declining every year since 2008 hy. In the four years from 2008 hy to 2011 hy the highest level was 4.87 m on August 15, 2008 and the lowest one was -6.9 m on March 02, 2012. The difference between the two reached 11.85 m, which is the maximum fluctuation during this period (Fig. 3).

Secondly, annual average groundwater level
Table 2  Average groundwater levels (Elv. m) by hydrological year, by rainy and dry seasons and by month in the dry season.

<table>
<thead>
<tr>
<th></th>
<th>2008</th>
<th>2009</th>
<th>2010</th>
<th>2011</th>
</tr>
</thead>
<tbody>
<tr>
<td>Hydr Year</td>
<td>0.30</td>
<td>-1.72</td>
<td>-4.47</td>
<td>-5.37</td>
</tr>
<tr>
<td>Rainy Season</td>
<td>1.05</td>
<td>-0.48</td>
<td>-4.19</td>
<td>-4.85</td>
</tr>
<tr>
<td>Dry Season</td>
<td>-0.50</td>
<td>-2.94</td>
<td>-4.75</td>
<td>-5.84</td>
</tr>
<tr>
<td>Nv</td>
<td>1.23</td>
<td>-2.48</td>
<td>-4.87</td>
<td>-4.97</td>
</tr>
<tr>
<td>Dc</td>
<td>0.02</td>
<td>-2.42</td>
<td>-4.90</td>
<td>-5.91</td>
</tr>
<tr>
<td>Jan</td>
<td>-0.68</td>
<td>-2.62</td>
<td>-4.73</td>
<td>-5.65</td>
</tr>
<tr>
<td>Fb</td>
<td>-0.55</td>
<td>-2.30</td>
<td>-4.40</td>
<td>-6.26</td>
</tr>
<tr>
<td>Mr</td>
<td>-1.09</td>
<td>-3.30</td>
<td>-4.88</td>
<td>-6.56</td>
</tr>
<tr>
<td>Ap</td>
<td>-1.53</td>
<td>-4.44</td>
<td>-4.70</td>
<td>*</td>
</tr>
</tbody>
</table>

*The data of April in 2011hy is unorganized.

was 0.30m in 2008hy, -1.72m in 2009hy, -4.47m in 2010hy and -5.37m in 2011hy (Table 2). Based on these numbers, the amount of decline in groundwater level was 2.02m between 2008hy and 2009hy, 2.75m between 2009hy and 2010hy and 0.96m between 2010hy and 2011hy, and the amount of decline has reached approximately 5.67m in the four years from 2008hy to 2011hy. This decline appears to be due to an increase in the amount of pumping. For instance, a pumping well G17, located 200m north of OW1, was installed in 2008 and began pumping in 2009. Wells in the PWG-west have been running since the second half of 2009.

Thirdly, the average value of the groundwater levels in the dry season of each hydrological year is lower than the value in the rainy season (Table 2). The groundwater level in the dry season is 1.55m lower than that in the rainy season in 2008hy, 2.46m in 2009hy, 0.56m in 2010hy and 0.99m in 2011hy. From these, it is evident that the lower groundwater level appears in the dry season of each year.

Fourthly, only in the dry season, the average value of the groundwater levels of each month is shown in Table 2. As shown in the table, the lowest groundwater levels appeared in April in 2008hy and 2009hy, in December in 2010hy and in March in 2011hy. As a trend, the critical decline of groundwater levels appears in the second half of the dry season, around March and April in particular.

In contrast, the groundwater levels at OW2 and OW3 have not changed as the level at OW1. In the time from 2008hy to 2011hy, OW2 records were in a range of approximately 2.5-3.0 Elv. m and OW3 showed a nearly stable level of -1.2 Elv. m. The reasons of these small fluctuations seem to be due to the distance from TLIP; the depth of screens and the hydraulic gradient towards the south. In this paper, these details are not considered.

V. Concluding remarks

Firstly, the following points are noted regarding the groundwater level at OW1. (1) The groundwater level has declined since 2008. The amount of decline
reached approximately 5.67m in between 2008/ly and 2011/ly. (2) The groundwater level appears lower in the dry season of each hydrological year, especially in March and April, and higher in the rainy season. (3) Why does the groundwater level rise during the rainy season? The reason seems to be that the water level of the Red River is rising in the rainy season and the infiltration and recharge from ground surfaces is increasing as well. Details will be analyzed in the future.

Secondly, the decline of groundwater level at OW1 is considered to be due to the large amount of pumping within TLIP. If this condition continues, the depletion of groundwater will eventuate in the near future. Companies that have expanded into TLIP should watch this situation carefully.

Thirdly, one of the urgent issues is to find methods to use the groundwater sustainably (Fischer et al., 2011; Hida, 2009; United Nations, 2012). An effective tool is to apply an artificial recharge of groundwater to the area of TLIP. In terms of quantity and quality, the water of the Ca Lo river is suitable as a source for the artificial recharge (Hida et al., 2010), and Kim Hoa Canal (Fig. 1) can be used as a recharge ditch.

Acknowledgement

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References


Hida, N. (2009): Managed aquifer recharge by using spreading basin methods on alluvial fans: a general overview of the situation in Japan. BOLETÍN GEOLÓGICO Y MINERO (special issue devoted to artificial recharge of groundwater), Instituto


HXBBD (2001): Topographic Map HA NOI and SON TAY (1:50,000), Hanoi Vietnam. (in Vietnamese)


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この論文に対する「討論」を2013年5月31日まで受け付けます。

ハノイ北西部のタンロン工業団地における地下水位の低下

肥田 登・Nguyen Van GIANG

要 旨

小論は、ベトナム、ハノイ北西部、タンロン工業団地における地下水位の低下について報告する。団地の日可能総揚水量は56,970 m³である。工業団地に隣接する観測井の記録によると、地下水位は2008年から低下し、2008年から2011年までの間に低下度はおおよそ5.67 mに達した。低下は年間を通じてみると乾季、特に3～4月に現れる。この団地の地下水利用を持続させるためには、Ca Lo川の水を引いている Kim Hoa水路を用いて地下水人工涵養を実施することが考えられる。

キーワード：地下水位、地下水人工涵養、タンロン工業団地、ハノイ